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Disentangling climatic and anthropogenic controls on global terrestrial evapotranspiration trends

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Abstract

We examined natural and anthropogenic controls on terrestrial evapotranspiration (ET) changes from 1982 to 2010 using multiple estimates from remote sensing-based datasets and process-oriented land surface models. A significant increasing trend of ET in each hemisphere was consistently revealed by observationally-constrained data and multi-model ensembles that considered historic natural and anthropogenic drivers. The climate impacts were simulated to determine the spatiotemporal variations in ET. Globally, rising CO₂ ranked second in these models after the predominant climatic influences, and yielded decreasing trends in canopy transpiration and ET, especially for tropical forests and high-latitude shrub land. Increasing nitrogen deposition slightly amplified global ET via enhanced

plant growth. Land-use-induced ET responses, albeit with substantial uncertainties across the factorial analysis, were minor globally, but pronounced locally, particularly over regions with intensive land-cover changes. Our study highlights the importance of employing multi-stream ET and ET-component estimates to quantify the strengthening anthropogenic fingerprint in the global hydrologic cycle.

1. Introduction

Intensified global hydrological cycle has been observed and modeled during the past few years (Huntington 2006, Gerten *et al* 2008, Wang *et al* 2010, Durack *et al* 2012, Douville *et al* 2013, Sterling *et al* 2013, Wu *et al* 2013, Gedney *et al* 2014). Terrestrial evapotranspiration (ET) is arguably the central component of this changing hydrologic cycle, and functions as a vital link between energy, water and carbon cycles, thereby having important implications for the availability and usage of fresh water resources by humans and terrestrial ecosystems (Seneviratne *et al* 2006, Trenberth *et al* 2009, Fisher *et al* 2011, Wang and Dickinson 2012).

Natural environmental factors (e.g. precipitation, temperature, incident solar radiation, soil moisture, wind and atmospheric teleconnections) regulate ET and its variability across different terrestrial ecosystems (Teuling *et al* 2009, Jung *et al* 2010, Wang *et al* 2010, Vinukollu *et al* 2011, Zhang *et al* 2012, Miralles *et al* 2014). These natural controls and limitations/co-limitations of ET are scale-dependent. Their mechanistic understanding is very important to predict the tendency and variability of ET (Wang and Dickinson 2012). Human-induced land use/land cover change, ground water withdrawals, and irrigation can directly alter the amount and timing of ET by modifying the local water and energy balances (Piao *et al* 2007, Gerten 2013, Leng *et al* 2013, 2014a, 2014b, Lo and Famiglietti 2013, Sterling *et al* 2013, Lei *et al* 2014c). Human activities that contribute to greenhouse gas emissions, atmospheric nitrogen deposition (NDE), and ozone pollution can also alter ET indirectly through changes in physiological, structural and compositional responses of plants (Gedney *et al* 2006, Betts *et al* 2007, Sitch *et al* 2007, Cao *et al* 2009, Leakey *et al* 2009). Discriminating these anthropogenic perturbations from natural factors is expected to increase in importance as anthropogenic transformation of the Earth System becomes more pervasive (Seneviratne *et al* 2010, Gerten 2013).

Based on mechanistic and empirical algorithms that are driven by remotely sensed observations, a variety of globally gridded diagnostic ET products have been compiled and evaluated in recent years (Willmott *et al* 1985, Fisher *et al* 2008, Jiménez *et al* 2009, Jung *et al* 2009, Sheffield *et al* 2010, Zhang *et al* 2010b, Miralles *et al* 2011, Mueller *et al* 2011, Vinukollu *et al* 2011, Zeng *et al* 2012, Schwalm

et al 2013). These gridded ET estimates offer crucial sources and benchmarks for quantitative investigations of historical ET dynamics over the land surface. However, the accuracy of these observation-based ET products has yet to be reconciled due to limitations in underlying hypotheses and errors in input datasets (Mueller *et al* 2011, 2013, Polhamus *et al* 2012). Moreover, due to their reliance on the satellite observations, these datasets offer a limited historical temporal record that encompasses only a few decades (Badgley *et al* 2015).

To predict future changes in ET patterns, process-based simulation and understanding of the magnitudes, mechanisms and interactions that control historical ET dynamics will be required and should be within uncertainty of both historical and present-day observations. Mechanistic land surface models (LSMs), driven by measurement-based environmental properties, are useful tools for the detection and attribution of natural and anthropogenic effects on ET dynamics. For the past decade, global factorial LSM experiments have been conducted and analyzed by different modeling groups to investigate the separate effects of environmental stresses on land surface and subsurface runoff, river flow, ET and water use efficiency (Gedney *et al* 2006, 2014, Piao *et al* 2007, Shi *et al* 2011, 2013, Tian *et al* 2011, Liu *et al* 2012, Tao *et al* 2014). The role of climate impacts on these hydrologic variables has been characterized predominantly across different regions of the globe. The relative role of natural environmental change versus anthropogenic activities, however, was modeled to be heterogeneous and geographically dependent. Nevertheless, due to large differences in initial model conditions, driver data, and complex parameterizations that govern models, the simulated ET was demonstrated to vary in magnitudes and responses across models at both temporal and spatial scales (Wang *et al* 2010).

To disentangle these differences in simulated ET patterns and the relative role of model sensitivity and structure, the experimental setup and boundary/initial data must be similar among different participating models. We leveraged the controlled factorial experiments and model simulation protocol from the Multi-Scale Synthesis and Terrestrial Model Intercomparison Project (MsTMIP) (Huntzinger *et al* 2013). Further, we synthesized a global ET time series (1982–2010) based on a diverse set of diagnostic ET products (table 1), and the methodology reported recently in Mueller *et al* (2013). The partitioning of ET

Table 1. Overview of the diagnostic ET datasets used for the merged ET of this study, and the simulated ET from MsTMIP models. Factorial results of the MsTMIP multi-model are ALL: the impact from all historical forcing factors, CLI: the impact from historical climate only, OTH: all anthropogenic impacts, CO₂: the historical CO₂ impact only, NDE: the historical nitrogen deposition impact only, LUC: the historical land use/land cover change impact only, Y: the availability of ET simulation for the particular impact, and N: the non-availability of ET simulation for the particular impact.

Group	Name	Algorithm	Spatial resolution	Precipitation data	Time period	Citation						
Diagnostic ET	GLEAM	Modified Priestley–Taylor	0.25° × 0.25°	GPCP CMORPH	1982–2010	Miralles <i>et al</i> (2011)						
	CSIRO	Modified Penman–Monteith	0.5° × 0.5°	SILO	1984–2005	Zhang <i>et al</i> (2010b)						
	MPI	Empirically derived from FLUXNET	0.5° × 0.5°	GPCC	1982–2008	Jung <i>et al</i> (2009)						
	NTSG	Modified Penman–Monteith	0.5° × 0.5°	GPCC	1983–2006	Zhang <i>et al</i> (2010a)						
	PRUNI (3 sets of data)	Penman–Monteith/Priestley–Taylor (ISCCP, AVHRR, SRB)	0.25° × 0.25°	Sheffield <i>et al</i> (2006)	1984–2007	Sheffield <i>et al</i> (2010)						
	PT-JPL	Modified Priestley–Taylor	0.5° × 0.5°	Not required	1984–2006	Fisher <i>et al</i> (2008)						
	UDEL	Modified Thornthwaite water budget	0.5° × 0.5°	GHCN2	1980–2008	Willmott <i>et al</i> (1985)						
	PUB	Empirical method (TWSA, CRU)	0.5° × 0.5°	GRACE	1982–2009	Zeng <i>et al</i> (2012)						
	AWB	Water balance	0.5° × 0.5°	GPCP	1990–2006	Mueller <i>et al</i> (2011)						
Group	Name	Algorithm	Spatial Resolution		Time Period	Citation	CLI	LUC	CO ₂	NDE	ALL	OTH
MsTMIP ET	CLM4	Modified Penman–Monteith	0.5° × 0.5°	CRUNCEP	1982–2010	Lawrence <i>et al</i> (2007), Mao <i>et al</i> (2012)	Y	Y	Y	Y	Y	Y
	DLEM	Penman–Monteith	0.5° × 0.5°	CRUNCEP	1982–2010	Tian <i>et al</i> (2011, 2012)	Y	Y	Y	Y	Y	Y
	BIOME-BGC	Penman–Monteith	0.5° × 0.5°	CRUNCEP	1982–2010	Thornton <i>et al</i> (2002)	Y	N	N	N	Y	Y
	CLASS-CTEM-N+	Modified Penman–Monteith	0.5° × 0.5°	CRUNCEP	1982–2010	Huang <i>et al</i> (2011), Bartlett <i>et al</i> (2006)	Y	Y	Y	Y	Y	Y
	CLM4-VIC	Modified Penman–Monteith	0.5° × 0.5°	CRUNCEP	1982–2010	Lei <i>et al</i> (2014a)	Y	Y	Y	Y	Y	Y
	ISAM	Modified Penman–Monteith	0.5° × 0.5°	CRUNCEP	1982–2010	Jain <i>et al</i> (1996)	Y	Y	Y	Y	Y	Y
	LPJ-WSL	Modified Penman–Monteith	0.5° × 0.5°	CRUNCEP	1982–2010	Sitch <i>et al</i> (2003)	Y	Y	Y	N	Y	Y
	ORCHIDEE-LSCE	Modified Penman–Monteith	0.5° × 0.5°	CRUNCEP	1982–2010	Krinner <i>et al</i> (2005)	Y	Y	Y	N	Y	Y
	SiB3-JPL	Penman–Monteith	0.5° × 0.5°	CRUNCEP	1982–2010	Baker <i>et al</i> (2008)	Y	Y	Y	N	Y	Y
	SiBCASA	Penman–Monteith	0.5° × 0.5°	CRUNCEP	1982–2010	Schaefer <i>et al</i> (2008, 2009)	Y	Y	Y	N	Y	Y
	TRIPLEX-GHG	Modified Penman–Monteith	0.5° × 0.5°	CRUNCEP	1982–2010	Peng <i>et al</i> (2011)	N	N	Y	Y	Y	N
	VEGAS	Bulk transfer formula	0.5° × 0.5°	CRUNCEP	1982–2010	Zeng <i>et al</i> (2005)	Y	Y	Y	N	Y	Y
	VISIT	Penman–Monteith	0.5° × 0.5°	CRUNCEP	1982–2010	Ito and Inatomi (2012)	Y	Y	Y	N	Y	Y

(e.g canopy transpiration (Tr) and evaporation from wet canopy and bare soil (ET–Tr)) and the variation of those ET components are poorly understood and less constrained by observations (Lawrence *et al* 2007, Jasechko *et al* 2013, Swenson and Lawrence 2014, Wang *et al* 2014). The MsTMIP modeling framework can advance our understanding of trends in ET by providing predictions of the individual ET components. In this study, we thus further investigated the contribution of individual influencing factors to the spatial and temporal characteristics of these ET constituents.

2. Datasets and methods

We created a merged diagnostic ET data (DIA) from 11 long-term diagnostic datasets, all based on different assumptions and constrained with extensive *in situ* observations or satellite retrievals or both (table 1). We remapped the monthly raw datasets from their original spatial resolutions to the half-degree resolution of the model output from 1982 through 2010 based on data availability. Following Mueller *et al* (2013), we applied both physical and statistical constraints for quality control and bias corrections. For the physical constraint, we developed a dataset of seasonal net radiation maxima using the surface radiation budget (SRB3.0) datasets (Gupta 1983). We then excluded grid points with values exceeding net radiation maxima by more than 25%. The outliers were identified as values that exceed ± 3 standard deviations (Weedon 2011). Then the median values of these quality-controlled multiple ET estimates were treated as the merged product, and were comprehensively compared with the LSM results in this study. As shown in figure S1, the annual anomalies of the previously synthesized ET in Mueller *et al* (2013) are well within the spread of this newly-merged diagnostic data product. This updated product however, provides longer-term dynamics and is more amenable for studies at multi-decadal timescales.

To isolate the contributions of environmental drivers to multi-year ET variations, we utilized the factorial ET simulations from the MsTMIP data archive. Driven by the same environmental forcing (climate variability and trends, rising atmospheric CO₂ concentrations causing fertilization and reducing stomatal opening, nitrogen deposition, land use/land cover change, and soil texture and vegetation types), these state-of-the-art LSMs were employed to identify the principal drivers of interannual variability and multi-decadal changes of ET. Because the evaporation component for canopy and soil, and the snow sublimation, were not separately archived in the standard model outputs in the MsTMIP I protocol (Huntzinger *et al* 2013, 2015), we included all relevant available outputs, namely the ET, Tr and the total evaporation (ET–Tr). Four model experiments: (1) SG1 (time

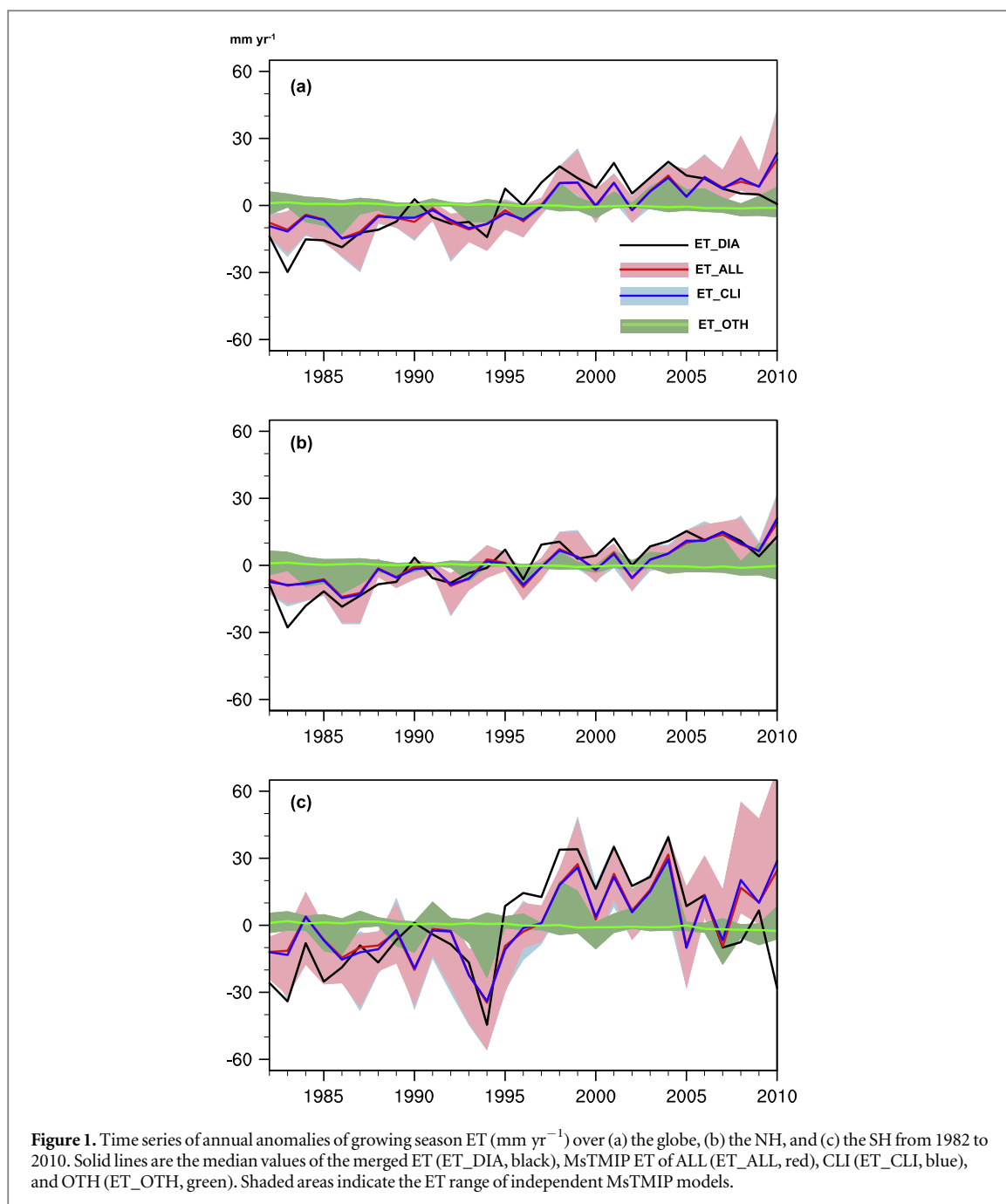
varying climate), (2) SG2 (time-varying climate and land use change history), (3) SG3 (time-varying climate, land use, and atmospheric CO₂), and (4) BG1 (time-varying climate, land-use, atmospheric CO₂ and nitrogen deposition), were analyzed to quantify the effects of each environmental forcing factor on the study variables for the years 1982 through 2010. The transient simulations began in 1901, turning on one time-varying driver at a time. Simulations BG1 or SG3 were used to address the combined impacts from various historical forcing agents for models with (BG1) or without (SG3) an explicit nitrogen cycle. Simulation or simulation differencing was used to quantify the contribution to ET and ET component changes from climate change (CLI) (derived from SG1), land use/land cover change (LUC) (derived from SG2–SG1), rising atmospheric CO₂ (CO₂) (derived from SG3–SG2), NDE (derived from BG1–SG3), or all forcing (ALL) (derived from BG1 or SG3) (table 1). To account for the overall effects from human activity (OTH), we derived the human-induced ET to be the difference between the BG1 and SG1 or SG3 and SG1 simulations.

Annual cropland area and total *tree* coverage information for the 1982–2010 period were derived from the merged product of the SYNergetic land cover MAP (Jung *et al* 2006) and the annual time series of the land use harmonization data (Hurtt *et al* 2011). Additional details on the aforementioned driver data and experimental design can be found in Wei *et al* (2014a, 2014b) and Huntzinger *et al* (2013, 2015).

Growing season ET generally dominates the annual sum over the vegetated area of land (Wang *et al* 2007). We focused our analysis on growing season ET for all observational and modeled data. The dynamic annual growing season information, used to mask the monthly ET between 1982 and 2010, was first determined from the global inventory modeling and mapping studies normalized difference vegetation index (NDVI3g) dataset (Pinzon and Tucker 2014) using a Savitzky–Golay filter (Chen *et al* 2004, Jonsson and Eklundh 2004). It was then refined by excluding the freeze period identified by the Freeze/Thaw Earth System Data Record (Kim *et al* 2011, 2012). In particular, the growing season of tropical rainforests was set to 12 months and it started in January.

3. Results

Across the globe, statistically significant increasing trends of ET were recorded from 1982 to 2010 in the observation-based ET estimates (DIA) (1.18 mm yr^{-2}) and modeled ET from the ALL simulation ($0.93 \pm 0.31 \text{ mm yr}^{-2}$) (figures 1 and S2, and table S1). Significantly positive annual correlations between the simulated ALL ET and the observed ET were obtained, particularly in the Northern hemisphere (NH) (Land: $R^2 = 0.58$, $p < 0.01$, NH: $R^2 = 0.72$,

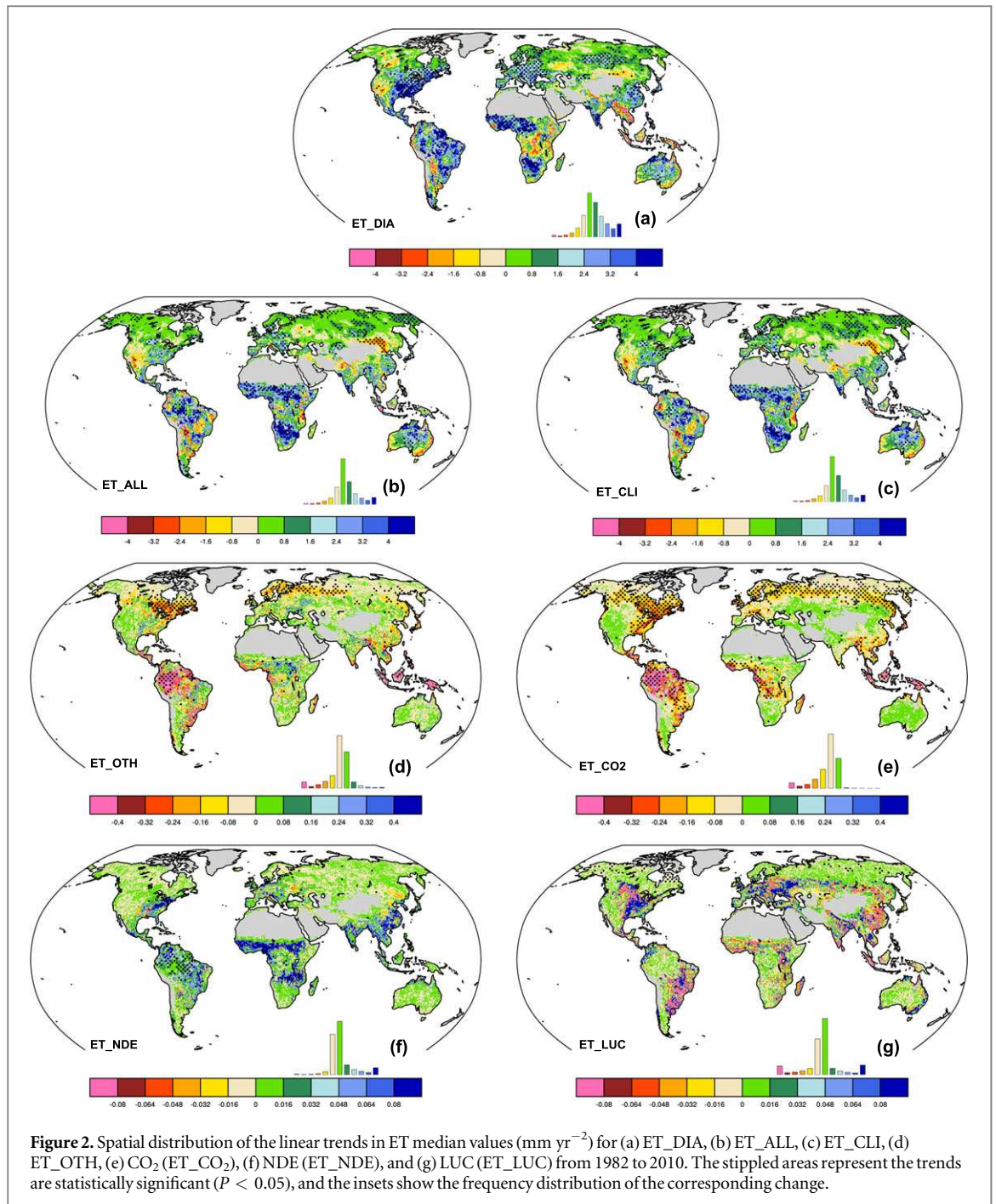


$p < 0.01$, and the Southern hemisphere (SH): $R^2 = 0.46$, $p < 0.01$). The simulated multiyear increasing trend and interannual variability of the ALL ET were mainly explained by the CLI ET. In contrast, the overall human-induced OTH ET was predicted to decrease somewhat, and to exhibit relatively small interannual variations.

Spatial analysis of linear trends of ET for the merged observation product revealed remarkably consistent increasing tendency over most continents (figure 2(a)). Local hotspots of reduced ET were diagnosed to occur in the arid regions of Western North America, central Africa, Northern China and South-eastern Asia. By contrast, the modeled changes of ALL ET underestimated the magnitude of ET changes in

Eastern North America and Western Europe, and missed the ET decreases in central Africa. But the placement of increasing or decreasing trends in ALL ET largely agreed favorably with those of the observed ET trends, indicating the suitability of examining multiyear ET trends using the all-factor simulations.

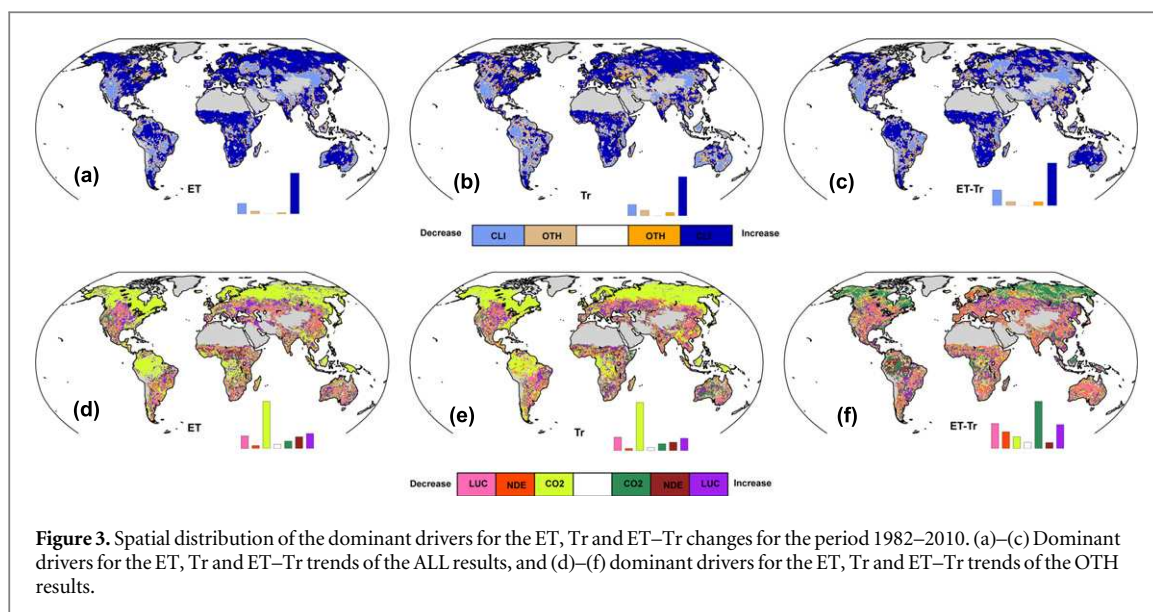
Spatial patterns of ET changes that are consistent between the ALL and CLM estimates confirm the dominance of climate forcing in explaining annual ET trends (figures 2(b), (c) and 3(a)). This dominant climatic response of ET trends was chiefly associated with concurrent annual precipitation changes (spatial $R^2 = 0.34$ for ALL ET and precipitation trend, and spatial $R^2 = 0.30$ for CLI ET and precipitation trends, respectively, $P < 0.01$), and tended to show large



spatial heterogeneities of sign and magnitude (figure 4(a)). The spatial dominance patterns based on the partial correlations among the total growing season ET, precipitation, temperature and incident solar radiation affirmed that for the MsTMIP models, annual precipitation drove not only the interannual variability of ET, but primarily accounted for the multiyear ET trends over most land areas (figures 4(e) and (f)). Combined anthropogenic effects tended to decrease ET, most notably in Northeastern North America, Western Amazon, Northwestern Europe and tropical Asia (figure 2(d)). These effects were subject primarily to the net physiological and structural

impacts of CO₂ concentration on the growth of plants in ecosystems (figures 2(e), 3(d), S2 and S3(a)).

Increasing nitrogen deposition led to increasing leaf area index (LAI) (figures 4(b) and S3(b)), and consequently to enhanced terrestrial ET, particularly over South America, Africa and Southeastern China (figures 2(f) and S2). The areas undergoing strong increase in forest fraction and decrease in cropland fraction, such as in central Eastern North America and central Europe, clearly showed increasing annual ET (figures 2(g), 4(c) and (d)). In contrast, regions with evident loss of trees, such as Eastern China and South-eastern South America, show a downtrend of annual



ET. Compared to the CO_2 and nitrogen deposition effects, however, the effect of LUC on land ET was important locally. Relatively large uncertainties from the LUC were also found between individual models (figures S2 and S6).

Trends for the Tr and total ET–Tr were dominated by the climatic changes across various continents. For Tr, 85.4% of the study area was impacted by the climatic changes, and 88.7% for ET–Tr (figures 3(b), (c), S4(a)–(f)). Congruent with the response of ET changes to rising CO_2 (48.4 ppm during the period 1982–2010), most areas, especially these regions covered by tropical broadleaf evergreen trees and high latitude shrubs, showed decreasing Tr. This is due to the CO_2 -induced reduction in stomatal conductance overwhelming the LAI-induced increase of canopy evaporation and transpiration under elevated CO_2 concentration (figures 3(e) and S4(j)). On the other hand, CO_2 fertilization would enhance canopy LAI through increasing photosynthate allocation to leaves, and caused more canopy transpiration and evaporation than the reduced transpiration by CO_2 physiological effects, especially over dry areas with sparse vegetation (e.g. the Western North America, central Eurasia, and Australia) (figures S3(a) and S4(j)). Reversed ET–Tr trends in these arid regions imply that decreasing soil evaporation was the dominant factor in changing ET–Tr (figures S4(j)–(l)). For most areas that showed decreasing Tr but increasing ET–Tr under CO_2 enrichment, the augmented evaporation of intercepted rainfall and increasing soil evaporation may have been coincidental.

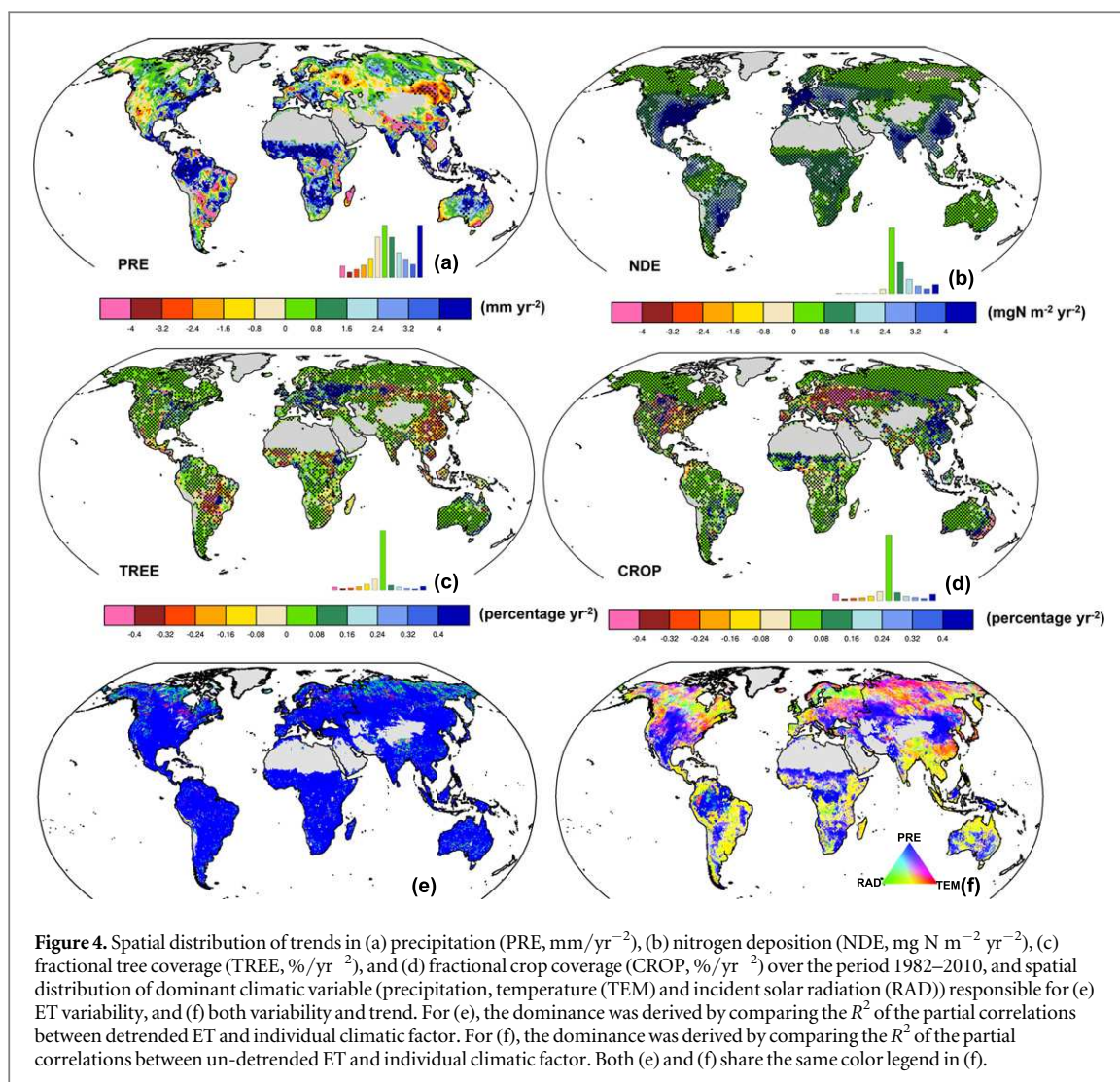
Increasing ET caused by nitrogen deposition was due to enhanced Tr (figures 2(f), S4(m) and S5). A decrease of ET–Tr caused by the nitrogen deposition effect, as seen in central North America and in Western Europe, was due to reduced soil evaporation (figures S4(n) and S5). The latter is a consequence of the increasing LAI providing more shade and so

reducing solar energy for soil evaporation. In addition, the increasing Tr further depleted soil water, which reduced soil evaporation. In the evergreen broadleaf forests of the Western Amazon and Congo basin, nitrogen deposition and higher LAI resulted in increasing canopy evaporation. The increase in canopy evaporation more than offset the decrease in soil evaporation and hence dominated the increasing ET–Tr and even the nitrogen-induced increase in total ET (figures S4(m)–(o)).

LUC led to a decreasing trend in Tr across densely inhabited regions that had experienced substantial land use perturbations (e.g. clearing trees for crops) during the study period. These occurred mainly in Southeastern South America and the Eastern China (figures 4(c), (d), S4(p) and S5). Tr trends showed a general negative sign over central Eastern North America and Western Europe, where croplands had been replaced mainly by forests and woodlands. This reduction of Tr with reforestation implies that the tree species that replaced the crops had lower stomatal conductance than the crop species, the younger and smaller trees of the returning forests had lower LAI than the croplands they replaced, or the available soil water for plants decreased because of the removal of irrigation. These aspects deserve further study.

4. Discussion

Between 1982 and 2010, the observation-based and simulated ALL ET consistently showed a significantly increasing trend across the globe. These findings are consistent with previous studies, which reported an intensified global hydrological cycle in response to global warming following the Clausius–Clapeyron law (the relationship between equilibrium water vapor pressure and temperature, about 7% per $^{\circ}\text{C}$ of warming) (Held and Soden 2006), as well as increasing



importance of the radiative component of ET (Johnson and Sharma 2010). Climatic factors accounted for much of the spatial and temporal variations in terrestrial ET, Tr and ET–Tr. This supports previous studies regarding the prevalent climatic mechanisms controlling the long-term ET trends such as temperature, precipitation, soil moisture, energy and internal climate variability (Teuling et al 2009, Jung et al 2010, Wang et al 2010, Vinukollu et al 2011, Zhang et al 2012, Ukkola and Prentice 2013, Miralles et al 2014).

In our study, the rising atmospheric CO_2 concentration, as tested by model factorial experiments, induced an overall suppression of Tr and hence a general decreasing ET. Our results further suggest that the sign of change and regional pattern of these CO_2 physiological effects on ET were moderated by changes in LAI. The overall response of ET was eventually determined by the balance among the changes of Tr, canopy evaporation and soil evaporation. These results are consistent with modeled and observed plant physiological responses to the increase of CO_2 concentration in the atmosphere (Betts et al 2007,

Leakey et al 2009). They also reiterate previous findings that show the concurrent physiological and structural responses of vegetation to rising CO_2 , and associated hydrological effects (Gedney et al 2006, Leipprand and Gerten 2006, Ainsworth and Rogers 2007, Betts et al 2007, Kurc and Small 2007, Piao et al 2007, Cao et al 2009, Leakey et al 2009, Lei et al 2014b).

Simulation experiments that consider NDE showed enhanced global LAI as a result of increasing nutrient availability (figures 4(b) and S3(b)). The nitrogen-induced enhancement of canopy Tr and canopy evaporation, however, was regionally offset by decreasing soil evaporation, and led to lower ET for the nitrogen fertilization effect. Nonetheless, mineralized nitrogen in the rooting system was governed by not only the amount of deposited N, but also by leaching and denitrification, which are affected by environmental conditions (Hovenden et al 2014). This highlights the necessity of better understanding the interactions among these environmental drivers, and the underlying mechanisms responsible for biogeochemical and hydrologic cycles.

Previous modeling studies (Boisier *et al* 2012, 2014, Shi *et al* 2013, Sterling *et al* 2013, Tao *et al* 2014) agree with our results that anthropogenic activities modified ET and its components locally, and human-induced LUC effects tended to counteract each other at a global scale. We found large uncertainties associated with LUC impacts among the MsTMIP LSMs, particularly over the NH and areas having marked land cover conversions. Though based on the same merged LUC dataset, different LSM groups prescribed the dynamic evolution of plant functional types with model-specific classifications (Wei *et al* 2014a, 2014b). The sensitivity of biophysical and biogeochemical processes to the reconstructed historical scenario of LUC, moreover, varied considerably from model to model (Huntzinger *et al* 2013). For example, for the SIB3-JPL models, abnormally higher LUC ET was simulated over the NH and global land compared to that of other models (figure S6). In SIB3-JPL, ET is a function of stomatal conductance and is sensitive to changes in photosynthetically active radiation (PAR). In LUC simulations, plant functional type changes over time, but the PAR is prescribed from present day NDVI climatology and is thus fixed to modern vegetation. This can lead to a bias in gross primary production in cases where grasslands are converted to forests, since the NDVI and resulting fraction of incident PAR absorbed by green leaving in the canopy ($fPAR$) are calculated from a modern day forest ecosystem but used to estimate stomatal conductance and ET for the historical grassland it replaced. The sensitivity to land-use change and cultivated ecosystems (e.g., irrigated croplands) reinforce the need for better LUC characterization, improved parameterization of ET in croplands, and the development of forcing datasets (e.g., PAR) that are not artificially dependent upon land cover. Improvements in these areas may help reduce the large inter-model spreads in the responses of ET to LUC.

Quantitative estimation of ET partitioning has been refined recently, but information on long-term variations and the precise drivers of each ET component are lacking (Jasechko *et al* 2013, Wang *et al* 2014). By using a multi-model ensemble, we assessed the annual trends of the Tr and ET–Tr over nearly three decades, and further estimated their spatial-temporal responses to various environmental stresses. These modeled results, however, remain rather uncertain without observational constraints that are sufficiently long and representative. Comprehensive synthesis of long-term observation-constrained ET components is needed to improve our understanding of the controlling mechanisms, and to better characterize the partitioning schemes.

5. Conclusions

The relative contribution of climate and anthropogenic activities to the spatio-temporal changes in ET

was quantitatively characterized with the newly-merged ET and multifactor ensemble simulations from MsTMIP. In the LSMs, climate, CO₂, nitrogen deposition, and land use impacts were separated experimentally to determine the ET variations between 1982 and 2010. Climate, and in particular, changes in precipitation, was the dominant control of multi-year ET trends and variability. The overall CO₂ physiological and structural effect induced decreasing plants transpiration and the total ET, especially in areas where vegetation was dense. Compared to climate change and the elevated CO₂ effects, the impacts of nitrogen deposition and land use change on ET were less important and acted locally. Other detailed explorations are needed, such as the implementation of more compelling statistical techniques and fully-coupled modeling systems (Douville *et al* 2013, Wu *et al* 2013, Gedney *et al* 2014) to detect and attribute the natural and anthropogenic effects on ET with more certainty. ET-related feedback studies are also required to account for land-atmosphere interactions and anthropogenic impacts in the integrated earth system models (Seneviratne *et al* 2010, Bond-Lamberty *et al* 2014, Collins *et al* 2015) and to understand future trajectories of drought (Sheffield *et al* 2012, Zarch *et al* 2015). Given that human activities continue to grow and intensify in the Anthropocene Epoch, we emphasize utilizing multi-stream datasets and multi-modeling frameworks to better diagnose and project anthropogenic influences on terrestrial ET, hydrologic cycle and overall climate change.

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